

Buoyancy Parameterization in Stress Generation

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- Stress is turbulence generated by wind shear and buoyance (density gradient).**
- Air is rarely exactly neutral in buoyance; even without temperature gradient there is always humidity gradient.**
- To get equivalent neutral wind or stress from wind, stability effect is needed.**
- Buoyance-induced bias may be small, but strong variations exist.**
- Stability effect was formulated 4 decades ago in land experiments; its geographical and seasonal validity over ocean has not been challenged.**

$$\tau = \rho C_D (U - U_s)^2$$

$$H = \rho c_p C_H (T - T_s)(U - U_s)$$

$$E = \rho C_E (Q - Q_s)(U - U_s)$$

$$U_* = \sqrt{\frac{\tau}{\rho}}$$

$$T_* = -\frac{H}{\rho U_*}$$

$$Q_* = -\frac{E}{\rho U_*}$$

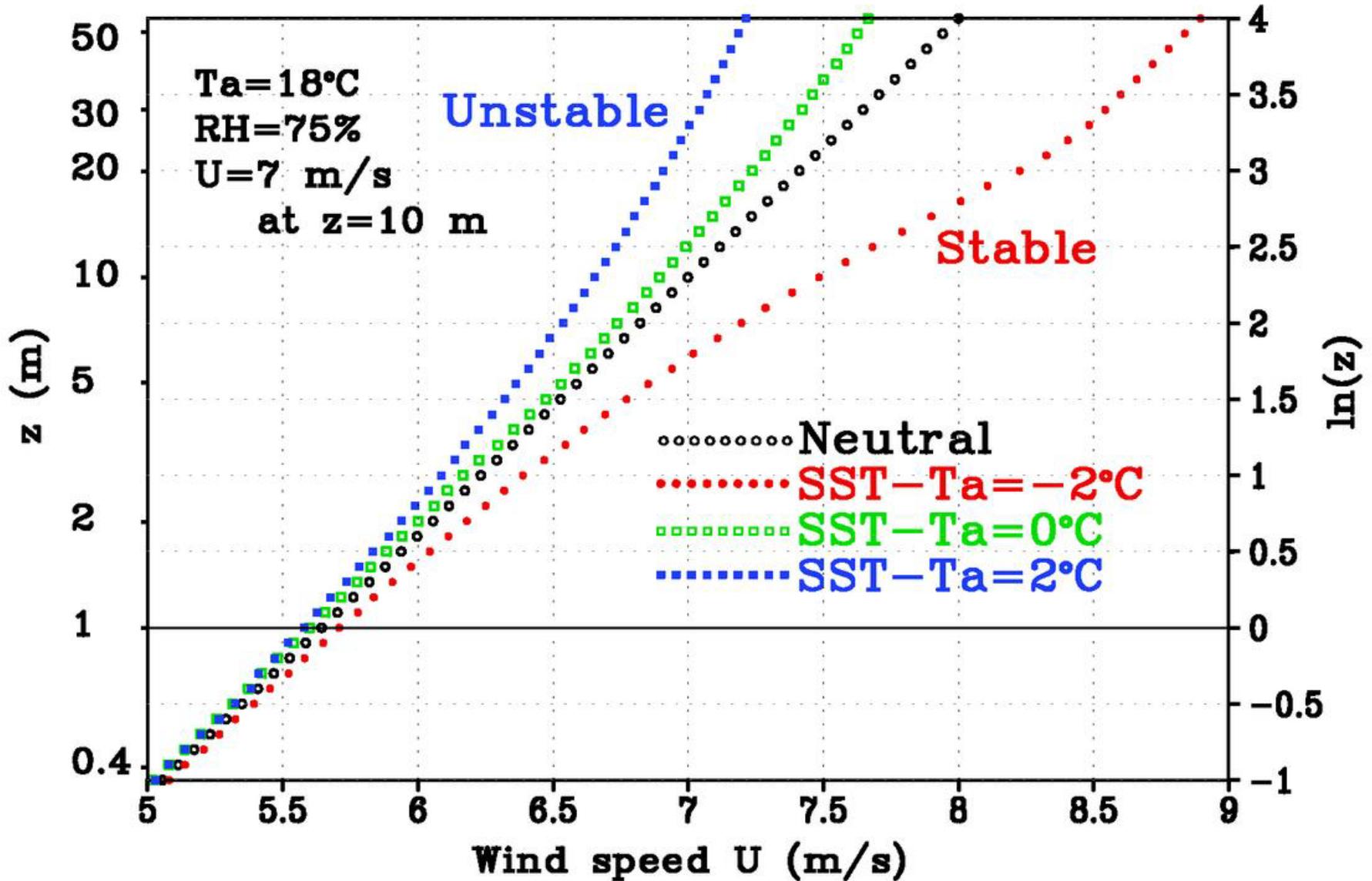
$$\frac{U - U_s}{U_*} = 2.5 \left(\ln \frac{z}{z_0} - \psi_U \right) = \frac{1}{\sqrt{C_D}}$$

$$\frac{T - T_s}{T_*} = 2.5 \left(\ln \frac{z}{z_T} - \psi_T \right) = \frac{\sqrt{C_D}}{C_H}$$

$$\frac{Q - Q_s}{Q_*} = 2.5 \left(\ln \frac{z}{z_Q} - \psi_Q \right) = \frac{\sqrt{C_D}}{C_E}$$

$$z_0 = 0.11 \frac{v}{U_*} + 0.011 \frac{U_*^2}{g}$$

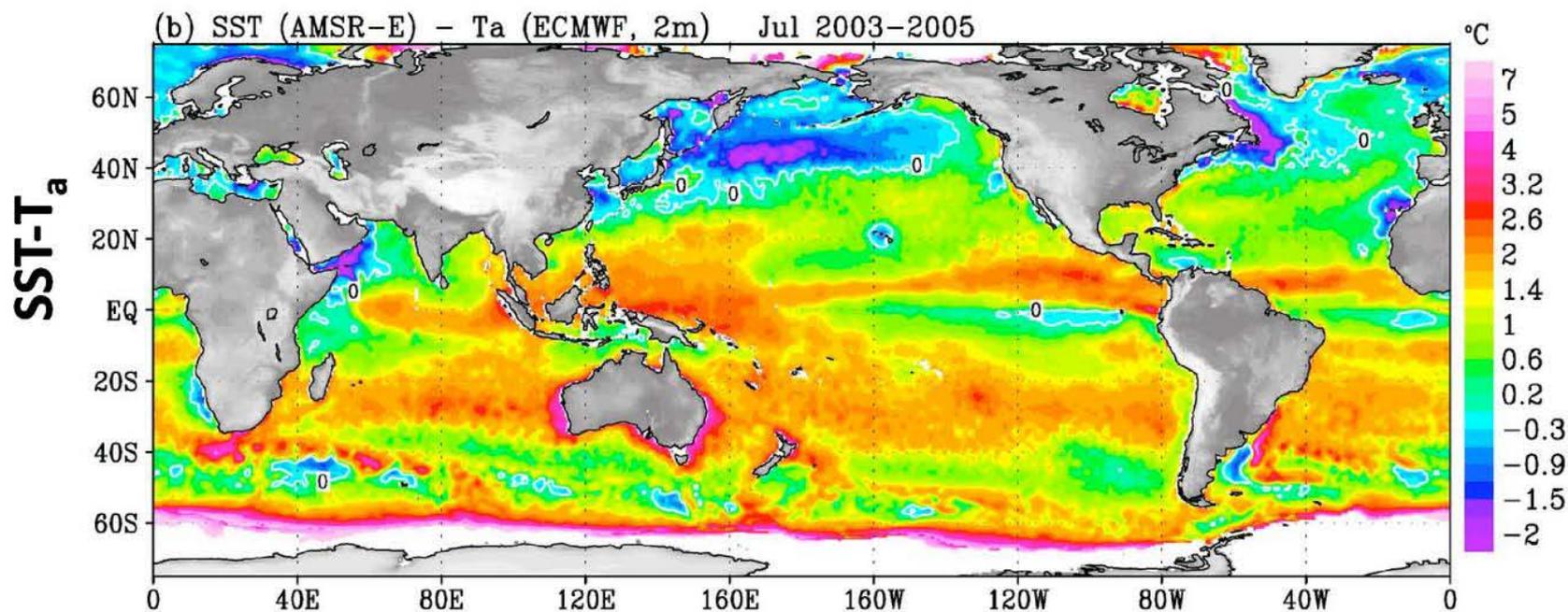
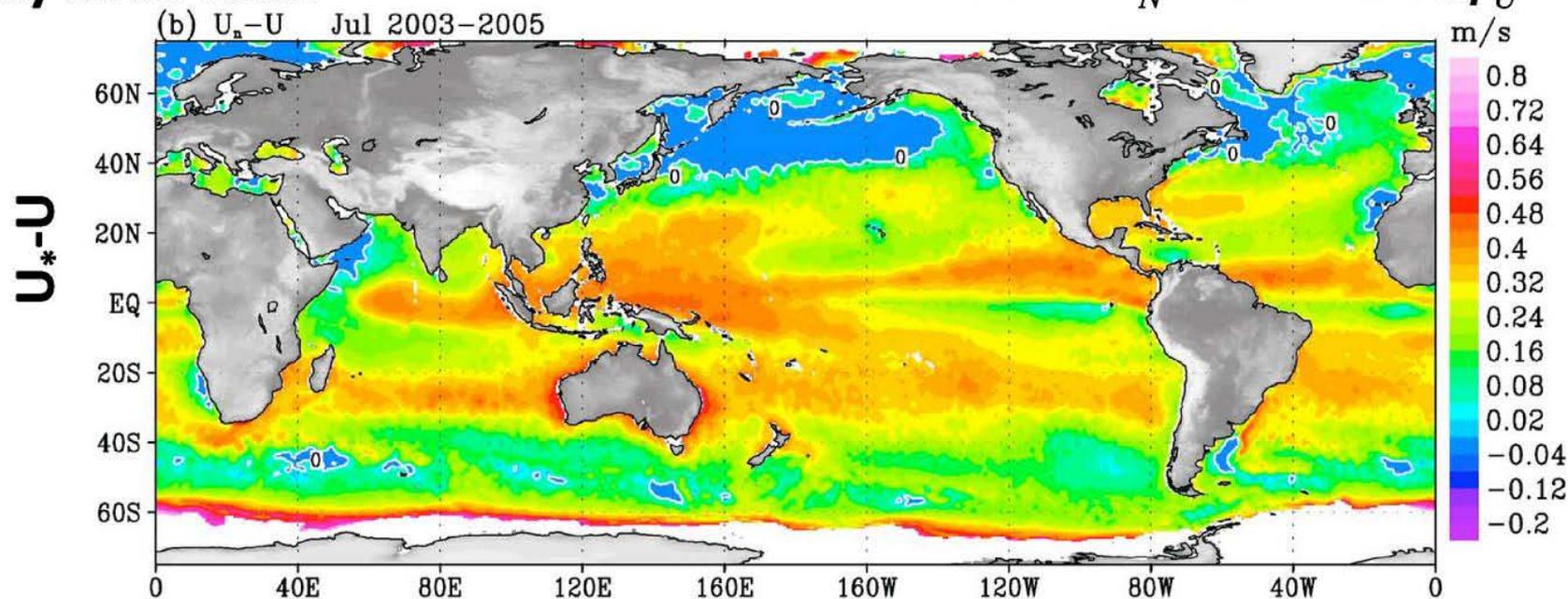
ENW is higher than actual wind under unstable condition



$$\frac{u - u_s}{u_*} = 2.5 \left(\ln \frac{z}{z_0} - \psi_u \right) = \frac{1}{\sqrt{C_D}}$$

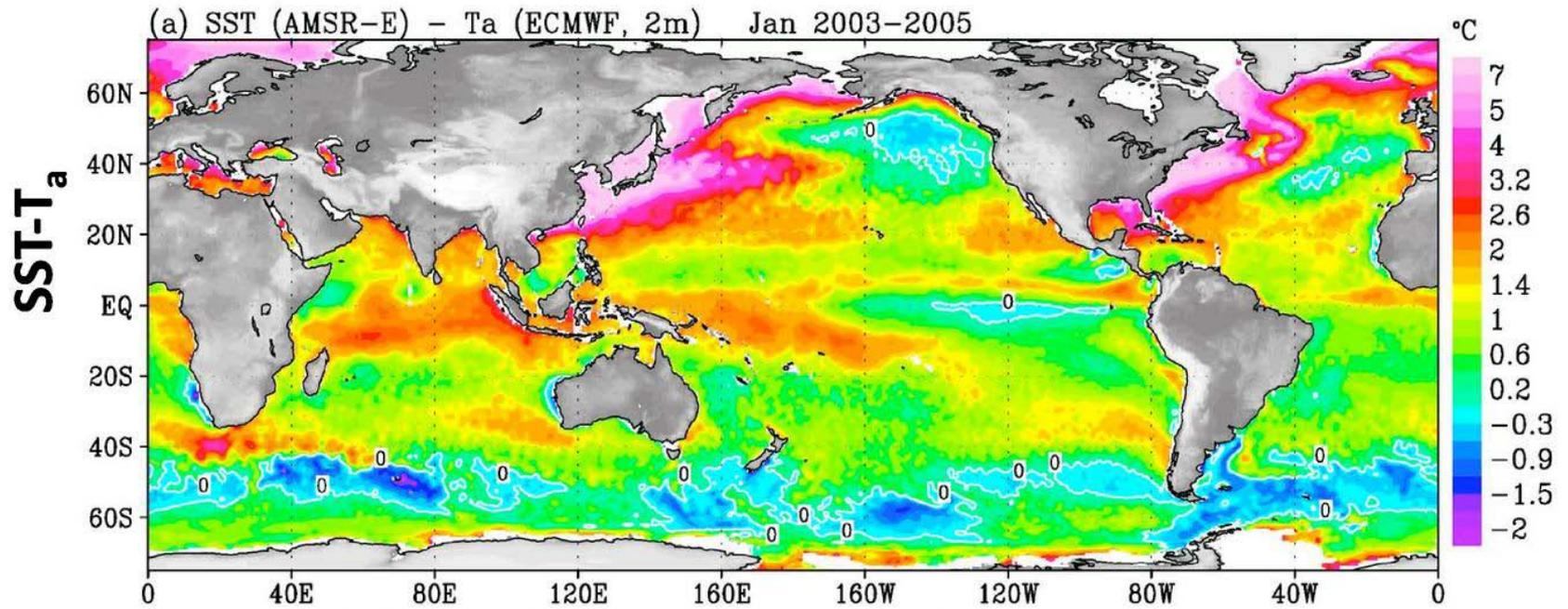
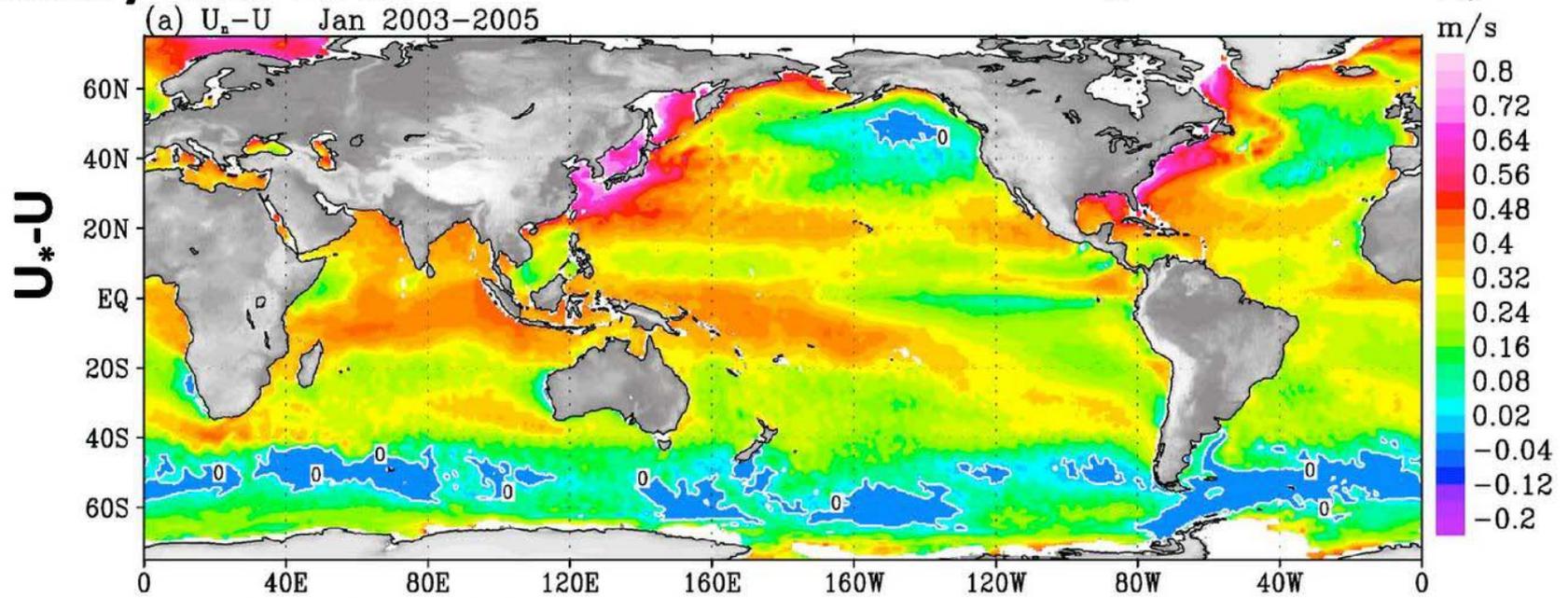
July 2003-2005

$$\delta U = U_N - U = 2.5U_*\psi_U$$



January 2003-2005

$$\delta U = U_N - U = 2.5U_*\psi_U$$



Similarity Theory

Buckingham theorem

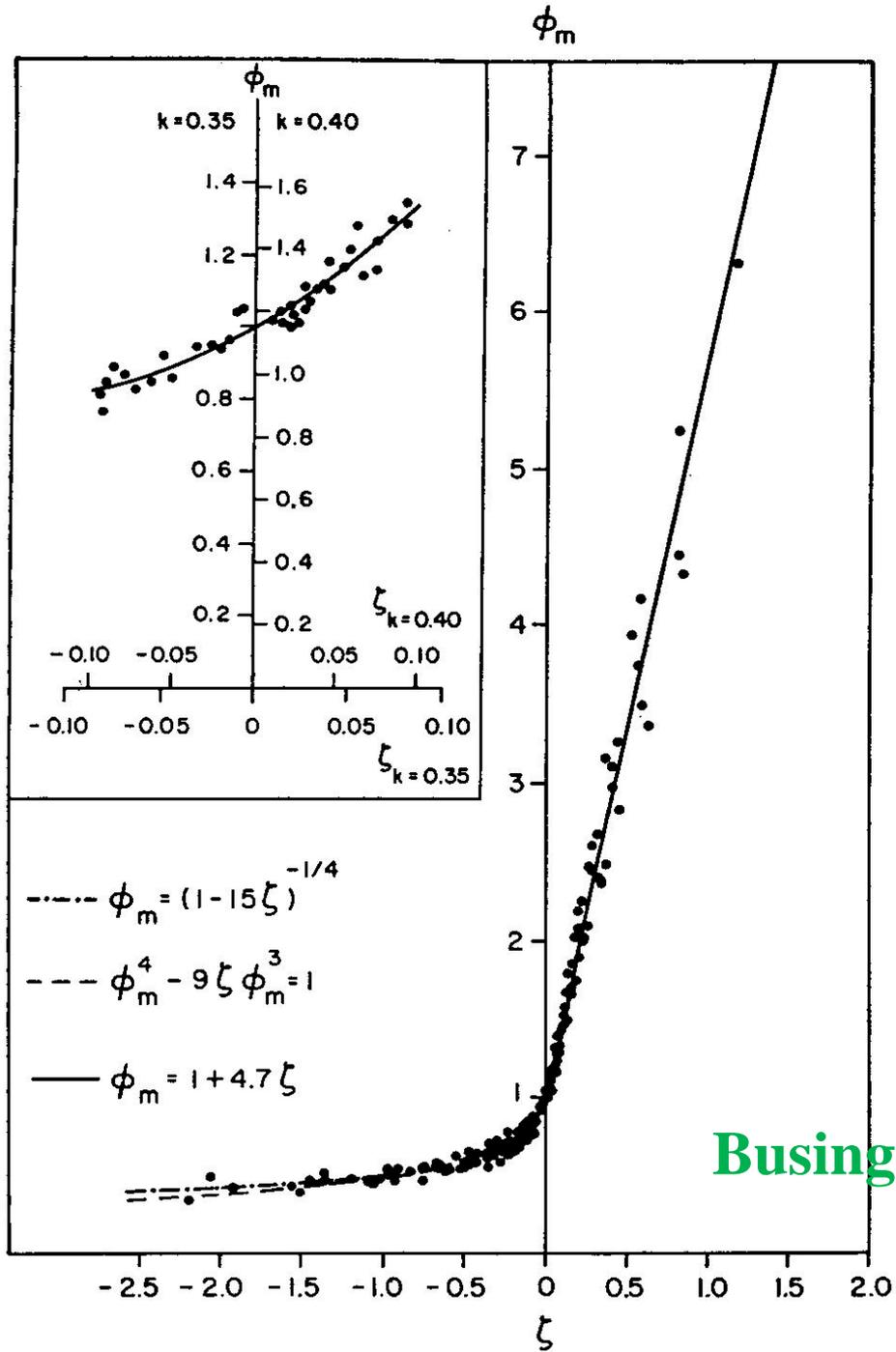
$$\frac{kz}{u_*} \frac{du}{dz} = \phi\left(\frac{z}{L}\right)$$

Obukhov (1946)

$$L = -\frac{u_*^3 c_p T}{kgH}$$

$$\phi = \left(-\gamma \frac{z}{L}\right)^{-\frac{1}{4}}$$

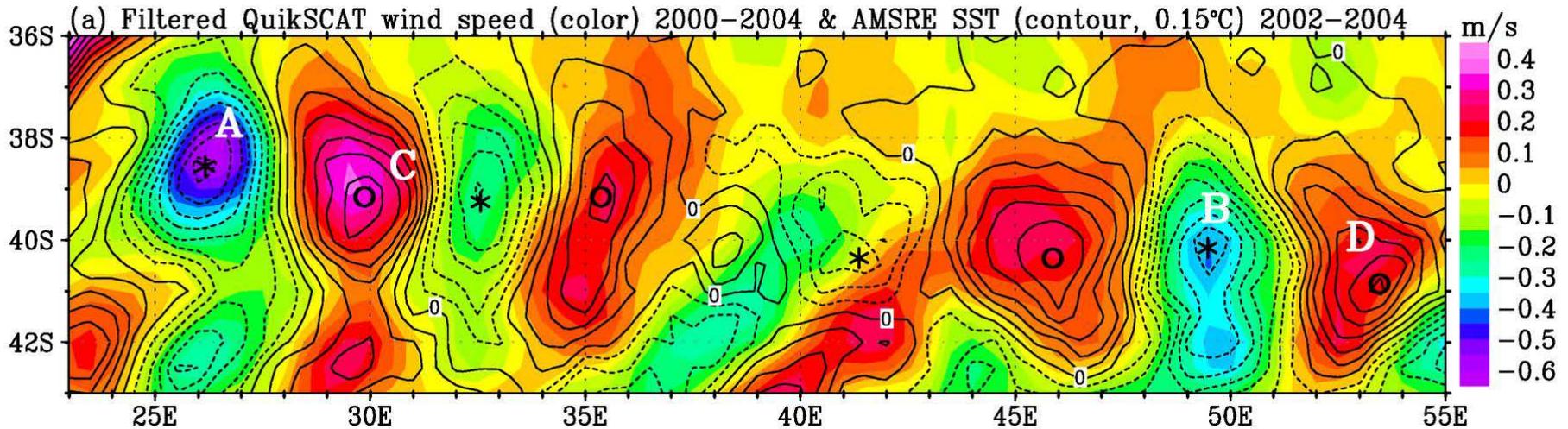
**Businger & Dyer model based on
Kansas & Kerang experiments**



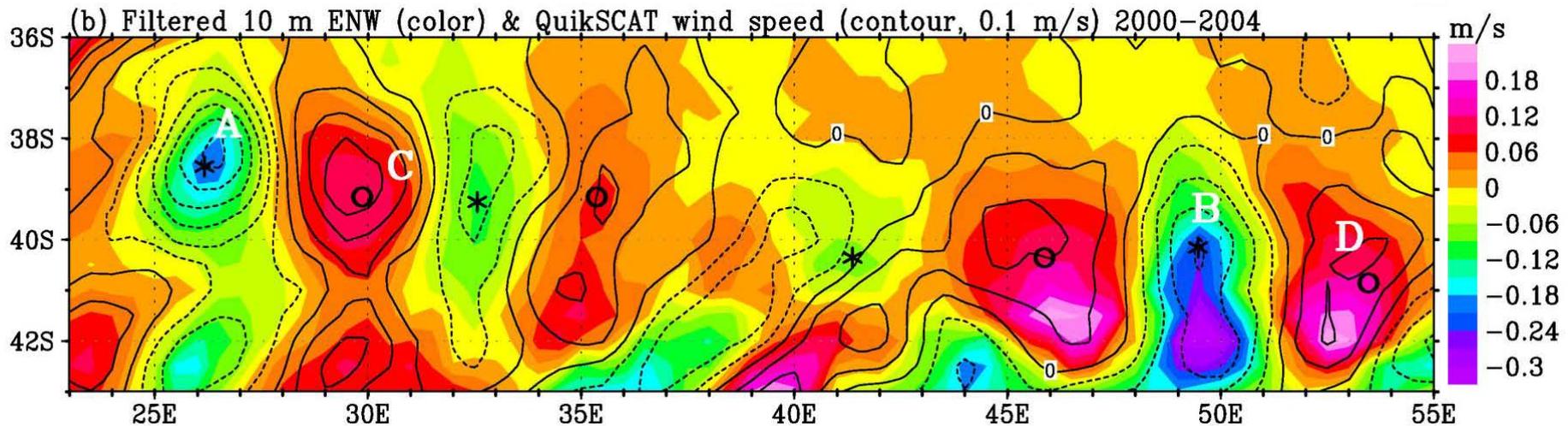
Businger et al (1971)

Observation from satellite

Agulhas



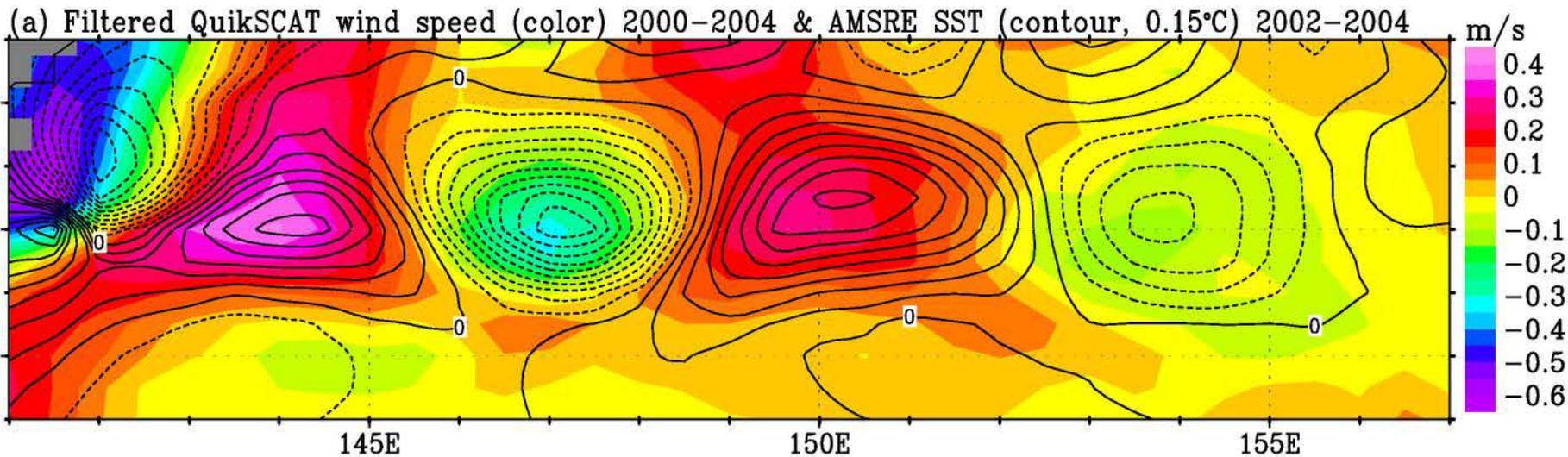
Computed from uniform wind field at 10m



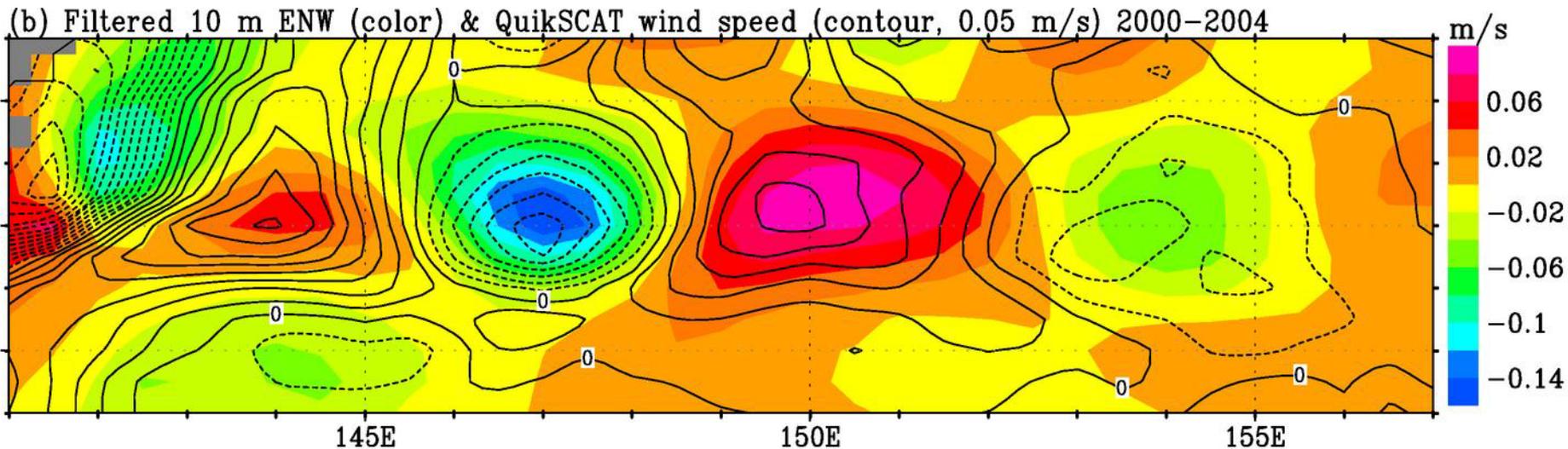
Collocation of ENW magnitude with SST is inherent in the definition of ENW and turbulent mixing theory. (Liu et al. 2007, JC)

Observation from satellite

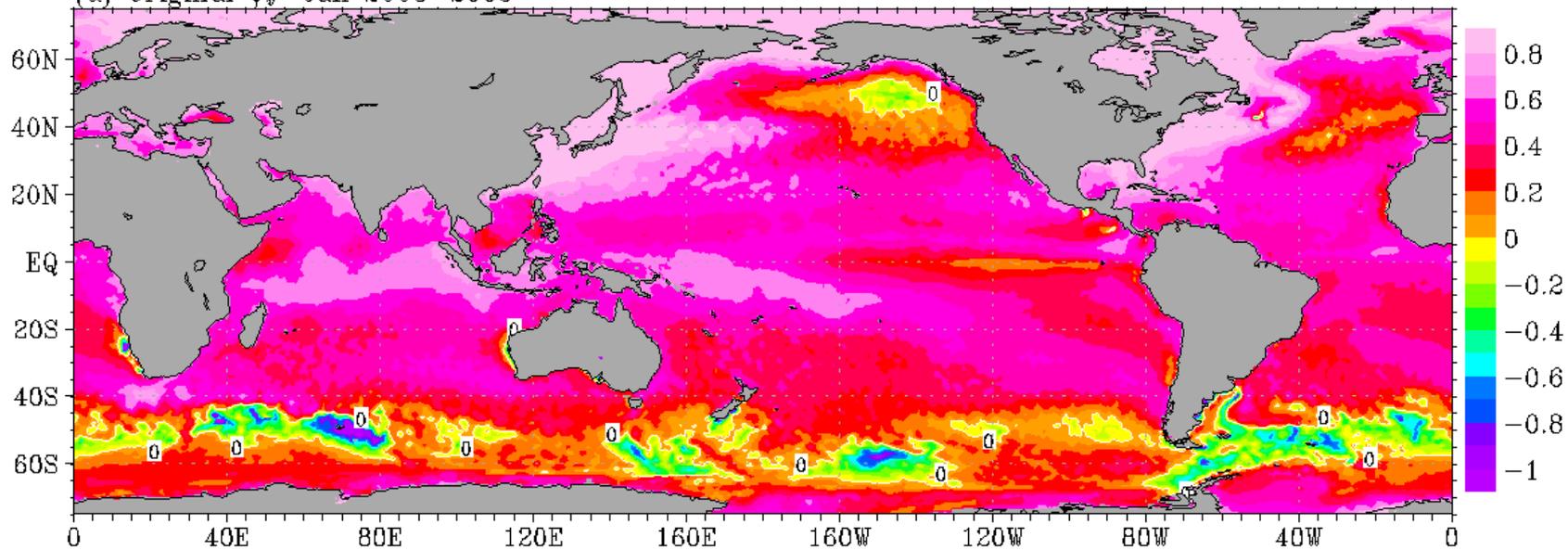
Kuroshio



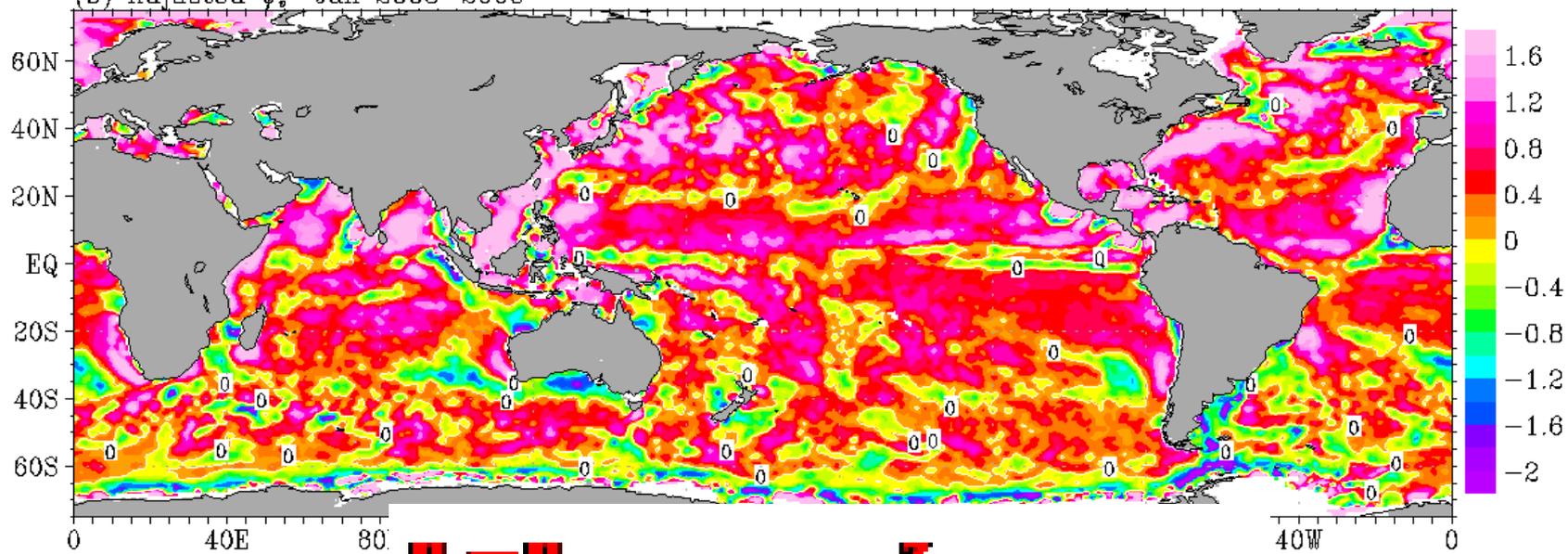
Computed from uniform wind field at 10m



(a) Original ψ , Jan 2003-2005

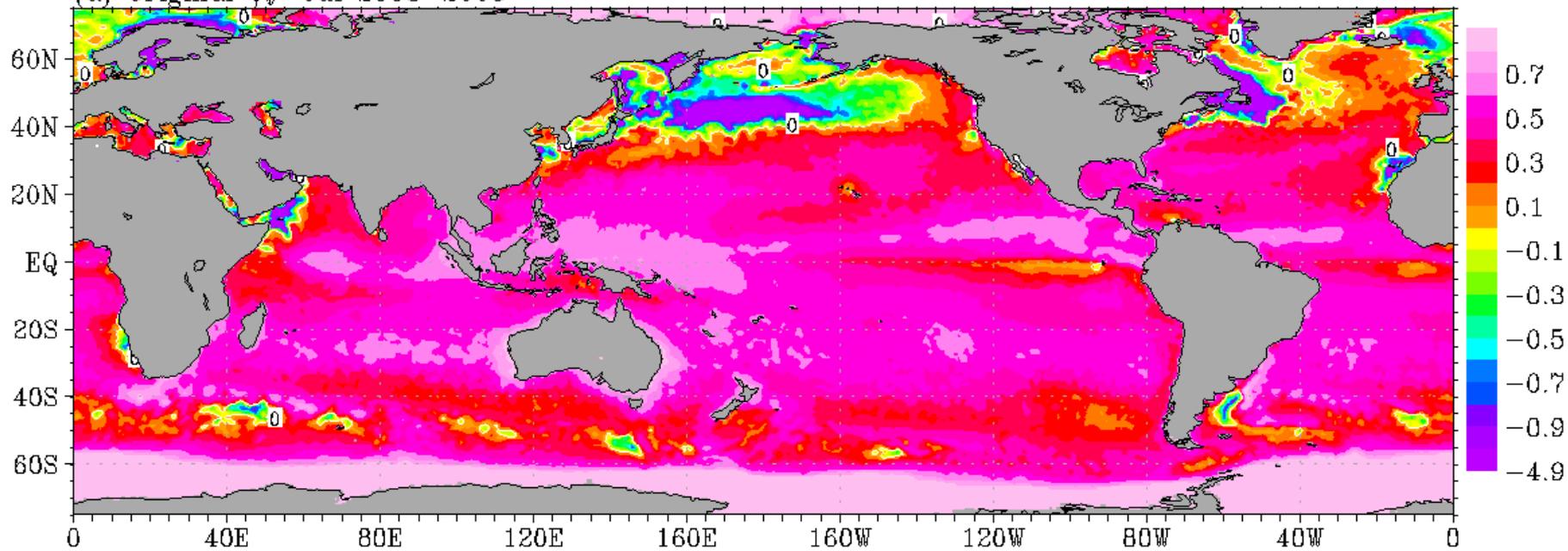


(b) Adjusted ψ , Jan 2003-2005

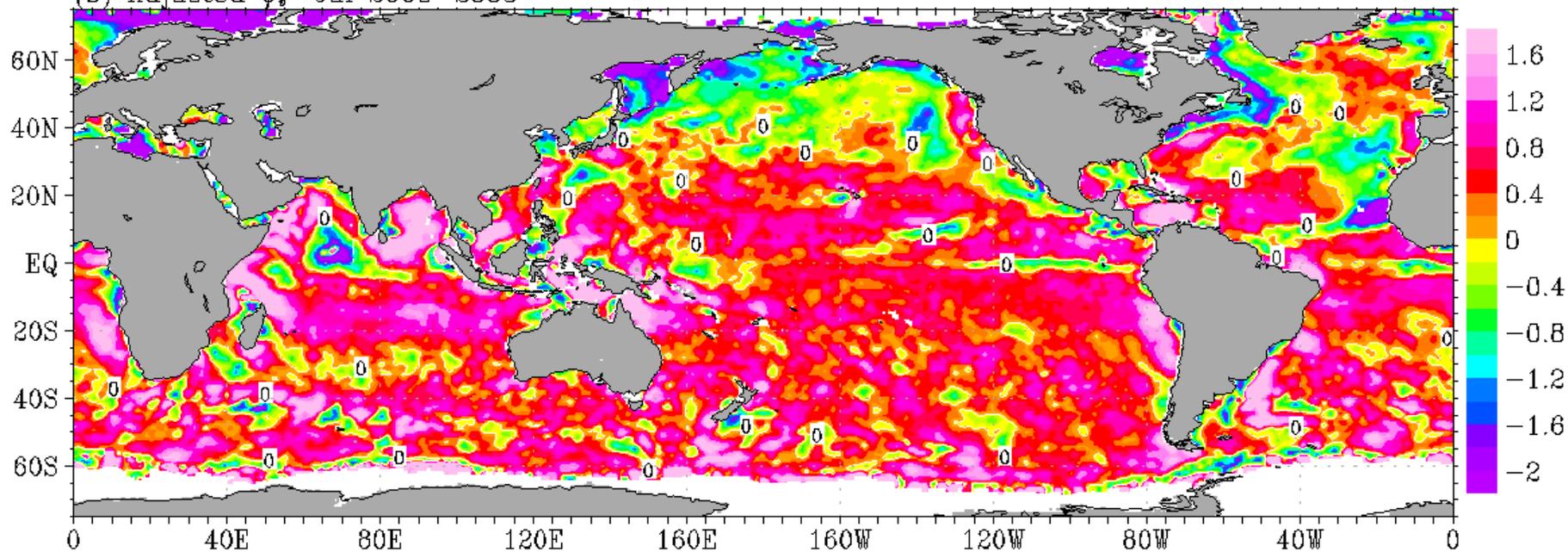


$$\frac{u - u_s}{u_s} = 2.5 \left(\ln \frac{z}{z_0} - \psi_n - \delta \right)$$

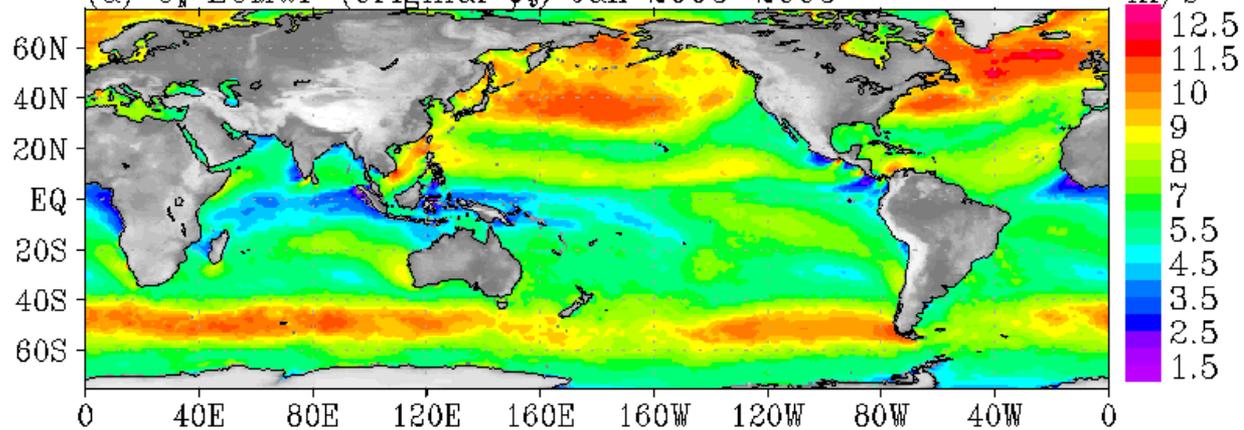
(a) Original ψ , Jul 2003-2005



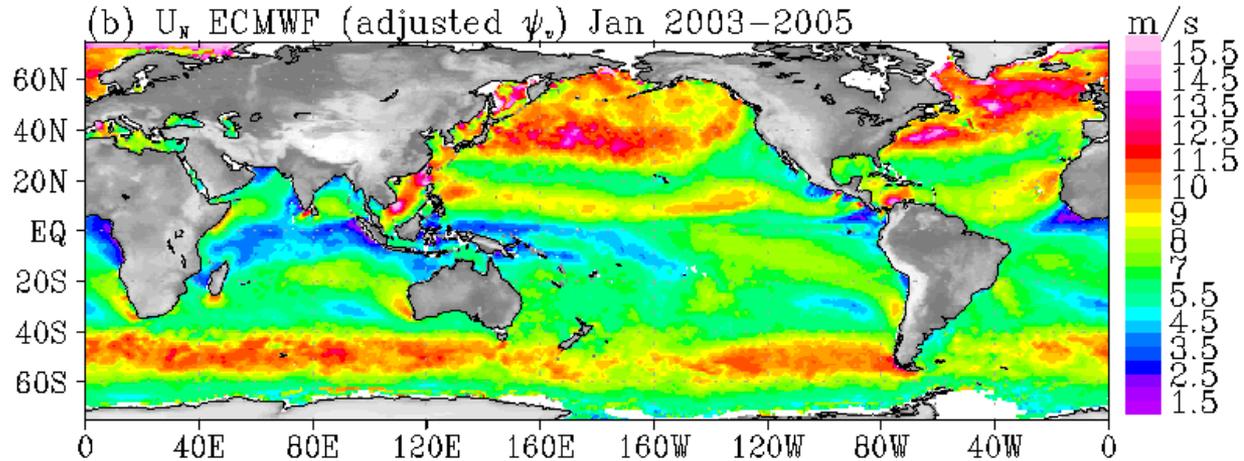
(b) Adjusted ψ , Jul 2003-2005



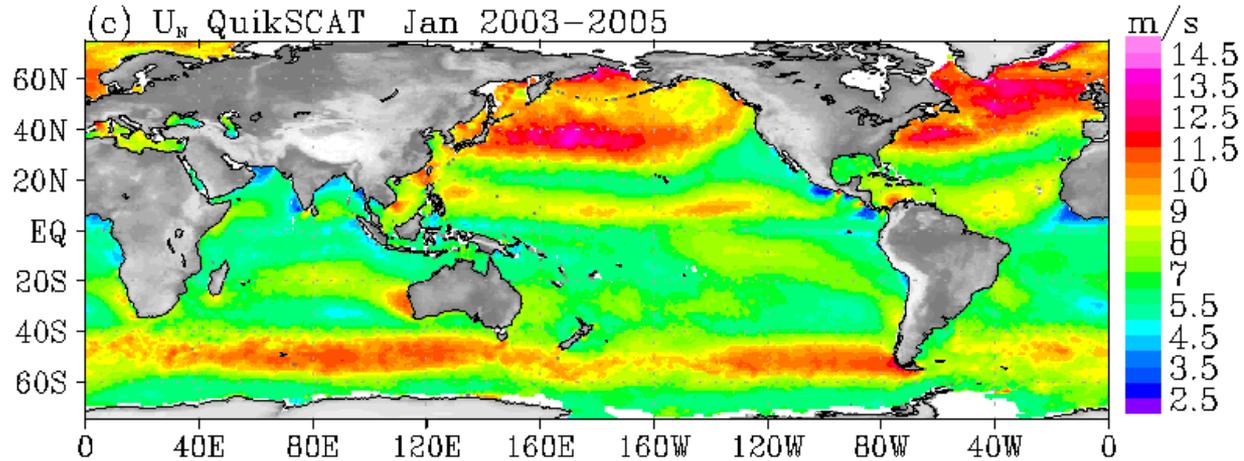
(a) U_N ECMWF (original ψ_v) Jan 2003–2005



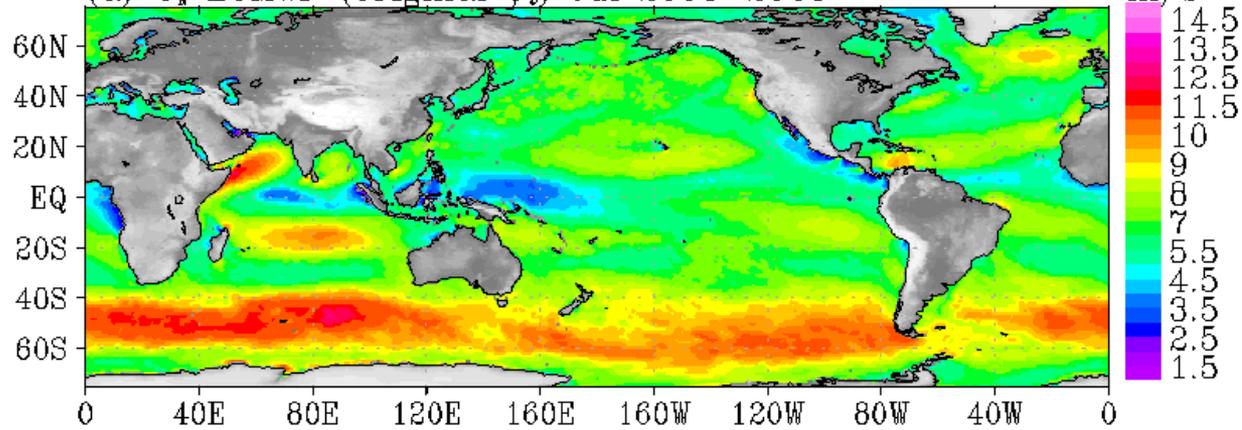
(b) U_N ECMWF (adjusted ψ_v) Jan 2003–2005



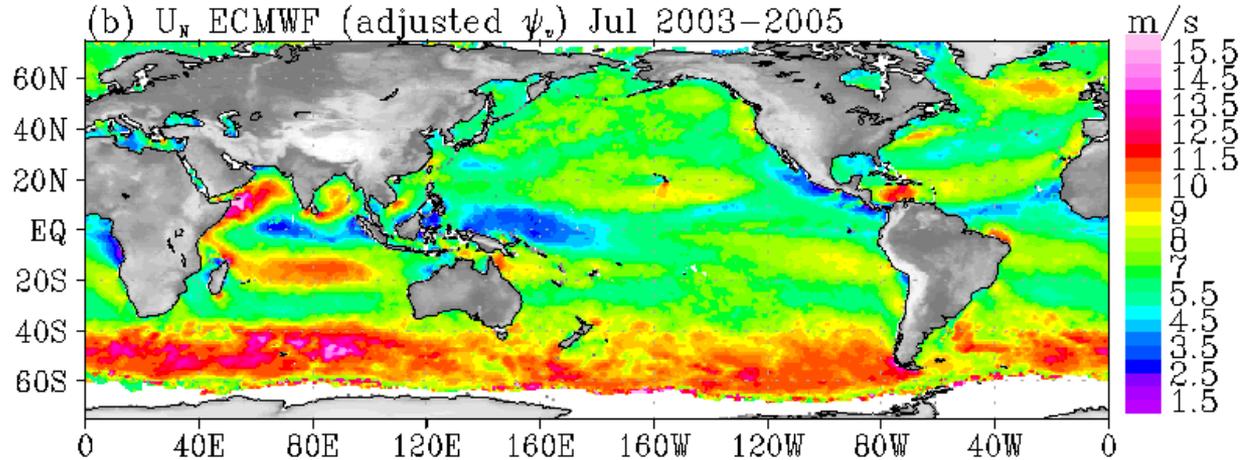
(c) U_N QuikSCAT Jan 2003–2005



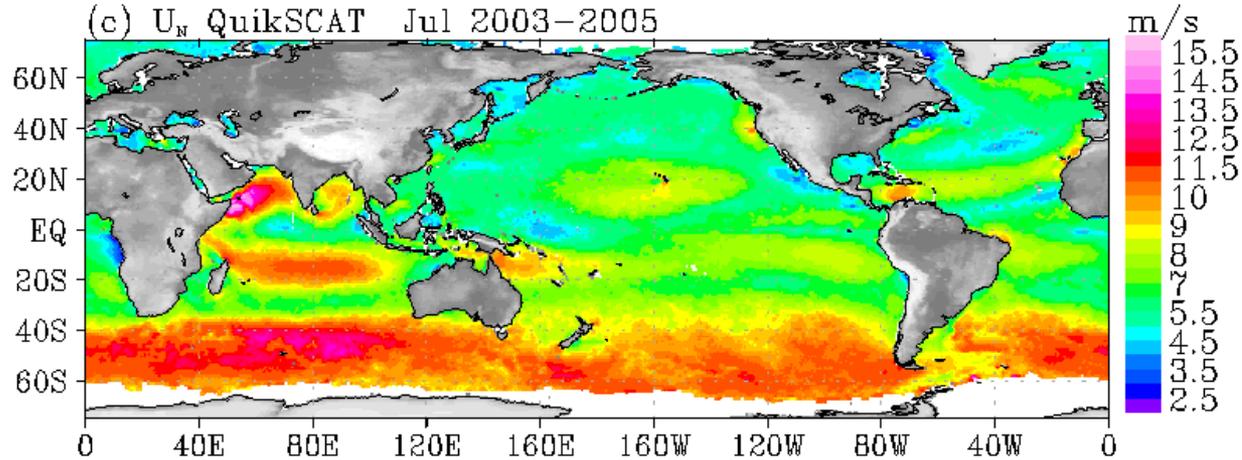
(a) U_N ECMWF (original ψ_v) Jul 2003–2005



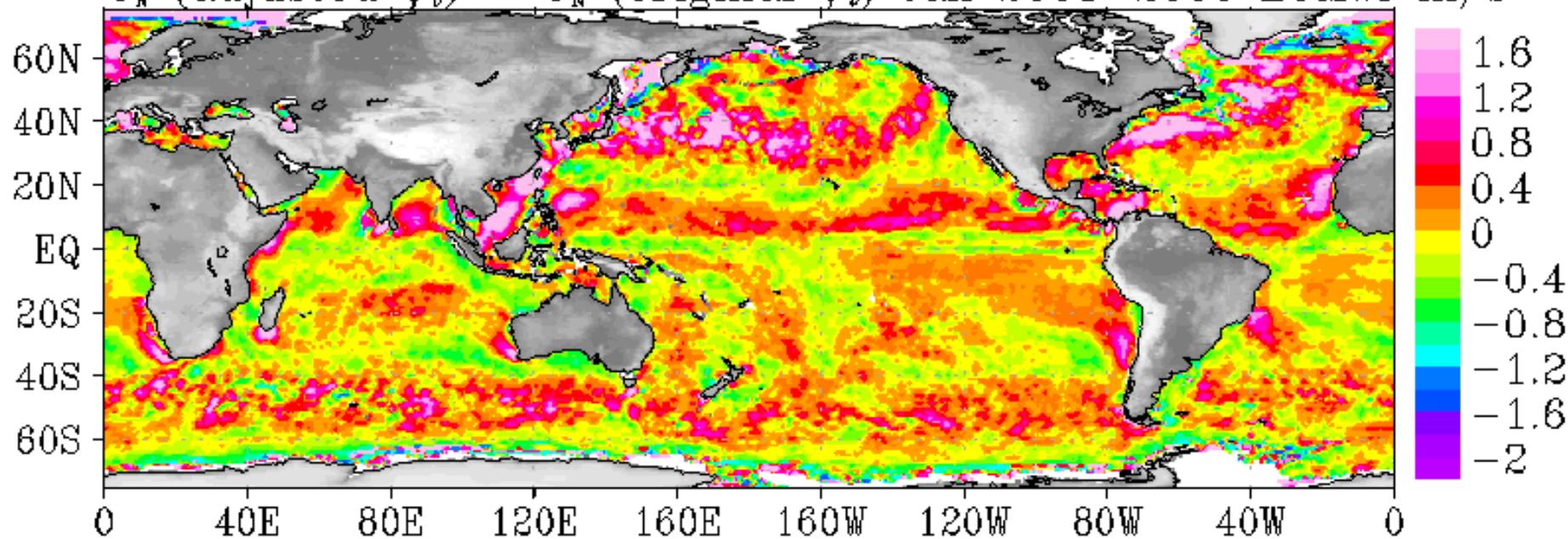
(b) U_N ECMWF (adjusted ψ_v) Jul 2003–2005



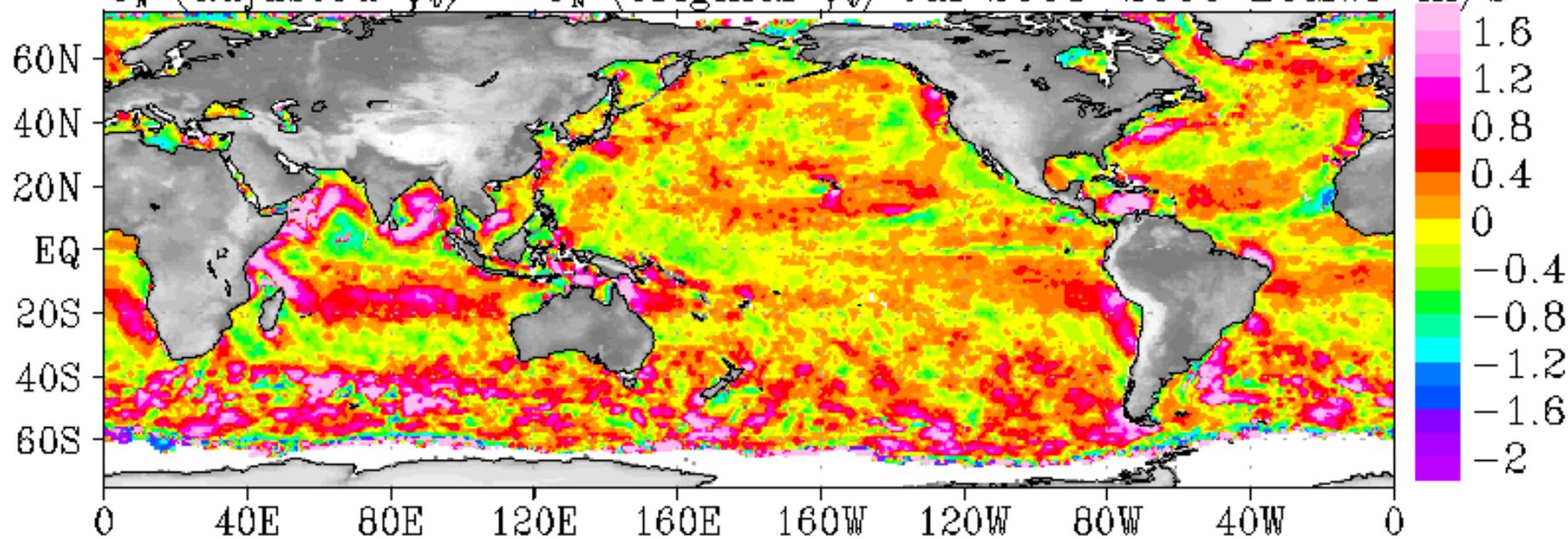
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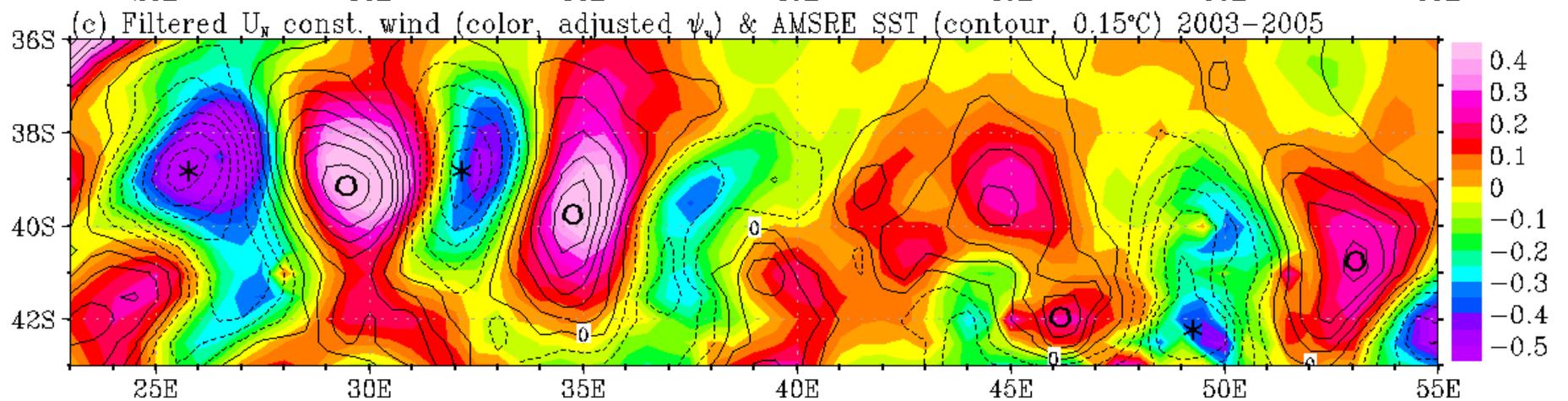
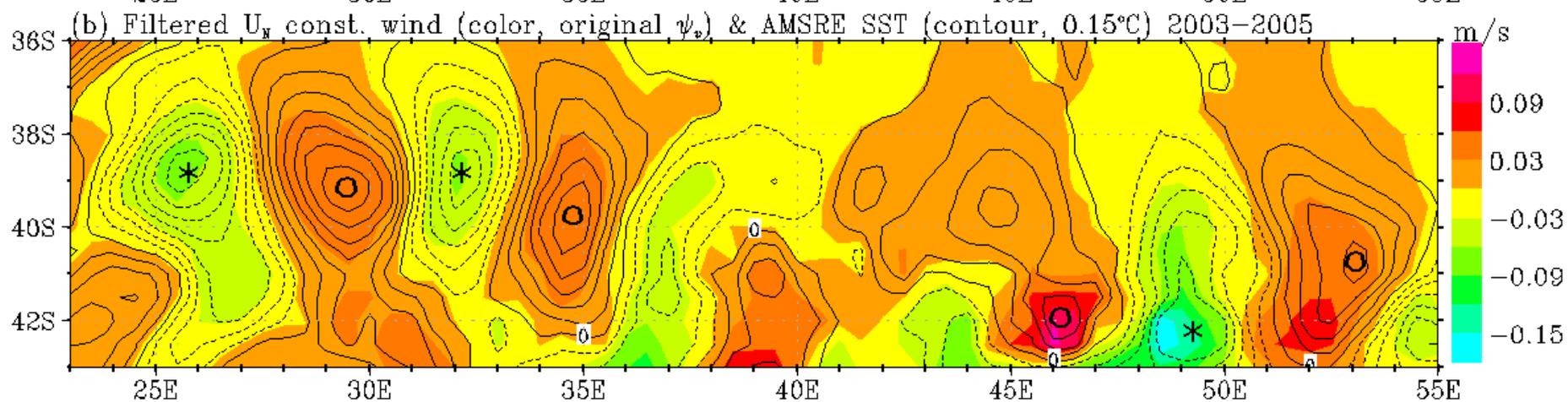
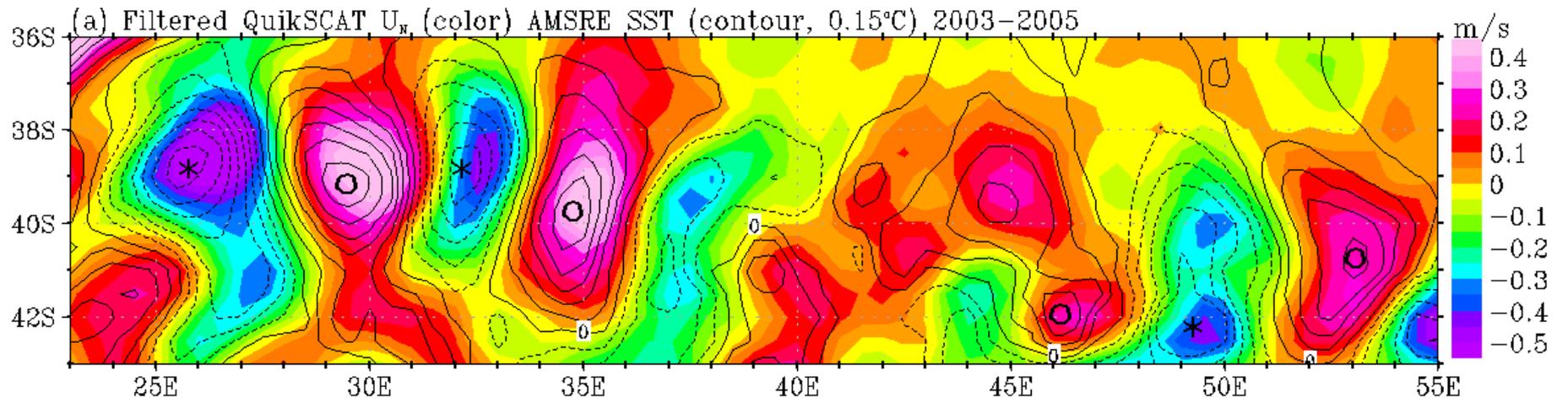


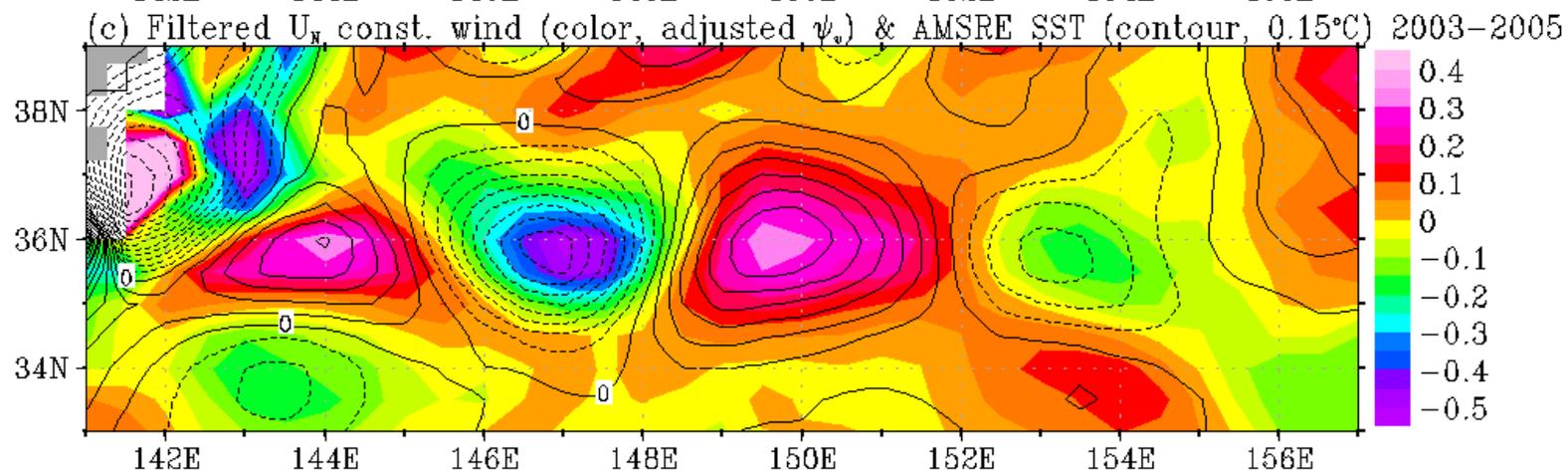
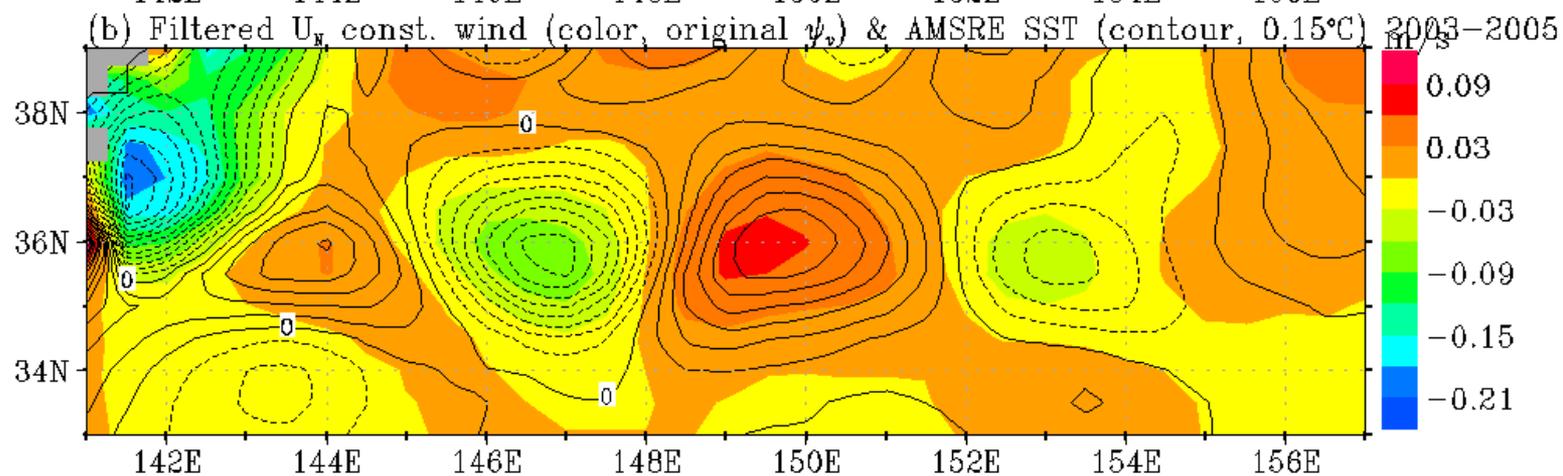
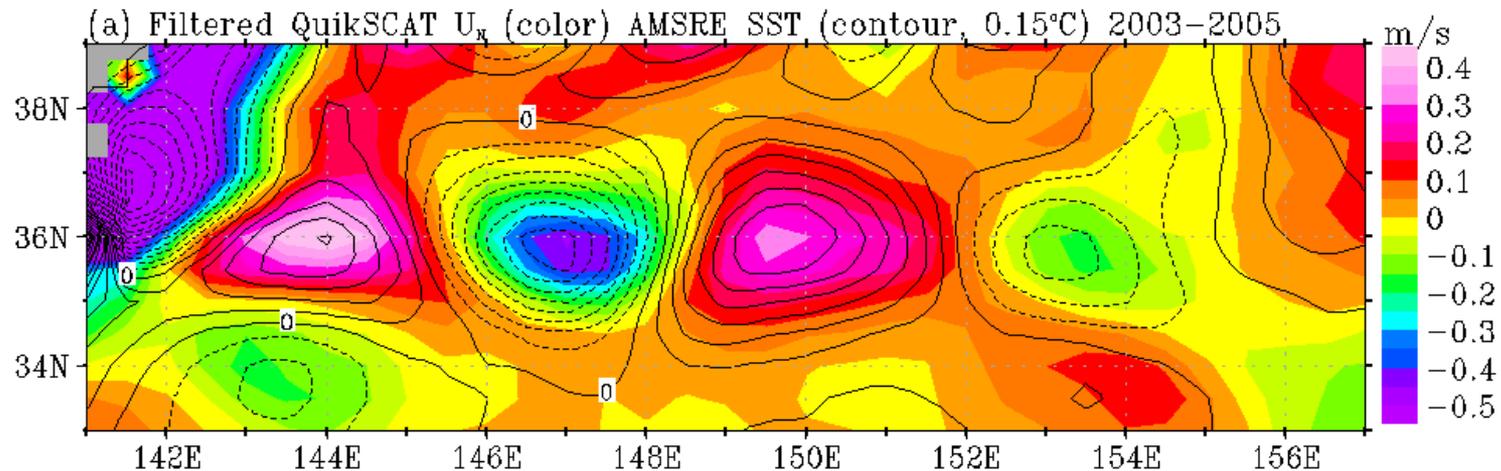
U_N (adjusted ψ_v) - U_N (original ψ_v) Jan 2003-2005 ECMWF m/s

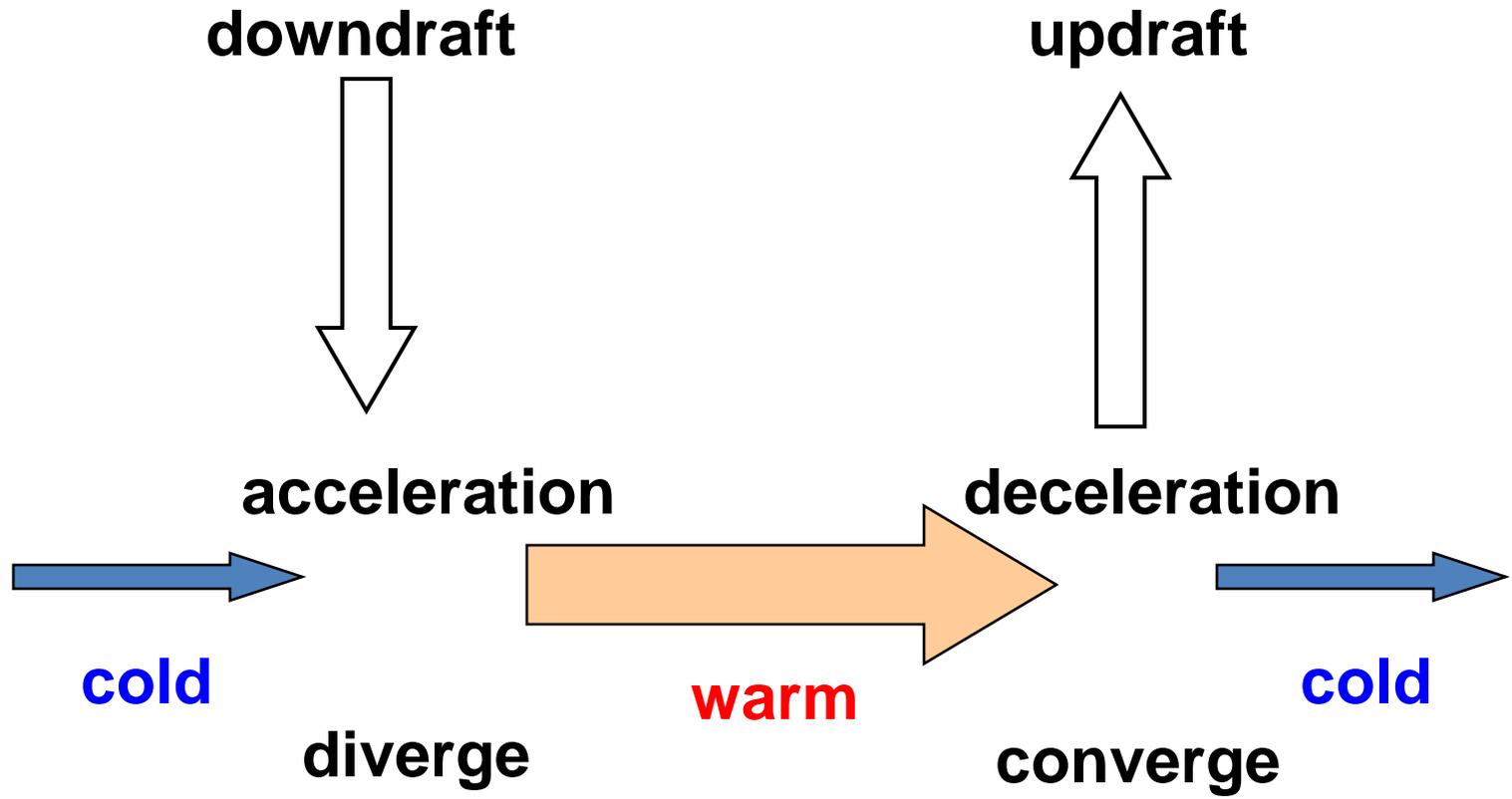


U_N (adjusted ψ_v) - U_N (original ψ_v) Jul 2003-2005 ECMWF m/s









- Our knowledge on stress came from wind through turbulence parameterization in the past
- Scatterpmeters provide the first and only maps of ocean surface stress
- The data generate new perspective of turbulence productions and its relation with mesoscale convection
- New bulk parameterization of stress may be needed

Backup